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3 **Calcite buffer effects in electrokinetic remediation of**
4 **clopyralid-polluted soils**

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22 **ABSTRACT**

23 This work presents a study of the dissolution-precipitation processes of calcite in soil
24 undergoing electrokinetic remediation processes. For this purpose, a numerical inspection
25 of one of these processes in a calcareous soil (10% calcite content) contaminated with an
26 of the organochlorine acid herbicide (3,6-dichloro-2-pyridinecarboxylic acid) was carried
27 out. The numerical tool used to perform this analysis was the Multiphysics for EKR
28 (M4EKR), a module programmed by the authors in the COMSOL Multiphysics platform.
29 A detailed analysis of the pH, the distribution of species that have a significant influence
30 on soil buffering capacity (carbonates and calcite), as well as the pollutant, has been
31 performed. In this way, the kinetics of the calcite dissolution-precipitation processes have
32 been analysed in order to determine the pH of the water in the pores and, therefore, how
33 it directly affects the effectiveness of the treatment. Additionally, a sensitivity analysis of
34 the selected EKR process has been carried out at different values of the rate of calcite
35 dissolution due to the uncertainty associated with the reactive surface area parameter.

36

37 **KEYWORDS**

38 Electrokinetic soil remediation, Multiphysics simulation, soil buffering capacity,
39 herbicide, M4EKR, calcite

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43 **1. INTRODUCTION**

44 The application of an electrical potential gradient between electrodes located in a soil
45 activates a series of electrokinetic transport mechanisms capable of mobilising water
46 (electroosmosis), ionic species (electromigration) and even charged particles
47 (electrophoresis). Specifically, this technique is called electrokinetic remediation (EKR)
48 when it is applied to contaminated soils with the aim of mobilizing polluting species to
49 controlled extraction points [1-4]. The coupling of the above-mentioned transport
50 mechanisms and the fact that, these treatments can be successfully applied in soils with
51 low hydraulic permeability, where Darcy flow is very limited, makes EKR technology a
52 very versatile technique and consequently, an efficient and competitive alternative for
53 treating pollution events with different compounds such as heavy metals [5-8], PAH [9-
54 13] or pesticides [14-18].

55 Regarding the EKR processes, the electrolysis of water is also an important
56 electrochemical reaction developed on the electrode surface. This process simultaneously
57 generates high concentrations of protons (H^+) by the oxidation reaction of water on the
58 anodic surface and hydroxyl ions (OH^-) in the cathodic reduction reaction. The changes
59 in the concentration of H^+ and OH^- are reflected in the generation of an important pH
60 gradient in the soil located between anodes and cathodes.

61 The pH changes generated in the vicinity of the electrodes depend on the amount of
62 electro-generated H^+ and OH^- and, in turn, on the electrical potential gradient applied in
63 the process. However, the pH-variations observed in the rest of the domain are determined
64 by the mobility of the acidic and alkaline fronts, from the anodic and cathodic zones
65 respectively. The transport of the pH front is mainly due to the electromigratory flow of
66 H^+ and OH^- species to the opposite sign electrodes and advective/diffusive processes, all
67 of which are strongly influenced by soil buffering capacity [19-24]. Generally, soil

68 buffering capacity is defined by the contribution of different processes: (i)
69 protonation/deprotonation of substances present in the porewater, (ii)
70 adsorption/desorption and ion exchange between the soil and the porewater and (iii)
71 dissolution-precipitation of minerals [25].

72 At this point it is important to stress the importance of pH changes in EKR treatments.
73 This variable directly influences the chemical speciation of the substances present in soil
74 and porewater, conditioning their ionic nature. Consequently, the pH of the system can
75 determine the predominant mechanism of transport of chemical species in the soil.

76 In order to have a global understanding of the evolution of pH in an EKR process, it is
77 necessary to improve the knowledge of the interactions attained between the
78 electrokinetic transport phenomena and the physical, chemical and electrochemical
79 processes that take place in an EKR treatment. The main objective of this work is to
80 evaluate the influence of soil buffering capacity on the performance of an EKR process.
81 For this purpose, a numerical analysis of the EKR treatment of a calcareous soil with a
82 calcite content of 10%, contaminated with clopyralid (3,6-dichloro-2-pyridinecarboxylic
83 acid, a common acid herbicide) was performed. Specifically, this study has evaluated the
84 effects generated by the calcite dissolution-precipitation processes on the general
85 behaviour of the EKR process and their importance for a better understanding of this type
86 of treatment.

87 The numerical tool used was the M4EKR module, Multiphysics for EKR [26]. M4EKR
88 is a numerical model developed by the authors as a result of the knowledge gained over
89 the last years in the treatment of contaminated soils using electrokinetic techniques [27-
90 35]. The M4EKR model has been implemented in COMSOL Multiphysics and is capable
91 of simulating the general behaviour trends of an EKR process by solving a coupled

92 problem of reactive transport and electric field. This makes the M4EKR module a
93 powerful computational tool that has been used to carry out different studies [36-38].

94 In this work, an assessment of the spatial distribution of the concentration of the species
95 that have a significant influence on soil buffering capacity (carbonates and calcite) was
96 performed. The relationship between the observed behaviour and the pH profiles obtained
97 is also analysed, as well as the direct influence on the performance of the simulated EKR
98 process. Additionally, and due to the difficulties in accurately predicting the kinetics of
99 calcite dissolution-precipitation, associated with the uncertainty of the reactive surface
100 area [39], a sensitivity analysis of the selected EKR process was performed using different
101 values of this parameter, which is equivalent to different levels of the rate of calcite
102 dissolution. Consequently, this study provides useful information on the influence of soil
103 buffer capacity on the general performance of an EKR process applied to the
104 decontamination of a calcareous soil with acidic contaminants.

105 **2. CONCEPTUAL MODEL AND NUMERICAL IMPLEMENTATION**

106 M4EKR is a reactive transport model applied to unsaturated porous media. It is a very
107 flexible numerical tool. Its adaptability and capacity for gradual implementation of
108 improvements are easy by the fact that the base module has been programmed in a
109 multiphysical environment. [26].

110 For conceptual simplification, the M4EKR version employed in this work does not
111 consider gas transport and the deformability of the soil. Due to the small scale of the
112 simulated tests (laboratory scale), it is acceptable to assume isothermal conditions (298
113 K). Given these simplifications, it is clear that the model will not perfectly match the
114 actual behaviour in an EKR process. However, this is not a limitation in the scope of the
115 work, since the aim is to improve phenomenological understanding and not to replicate a

116 real case. Soil buffering capacity is provided by: chemical reactions with proton
117 consumption/production (see Table 4), and dissolution-precipitation of calcite [40]. The
118 adsorption/desorption processes onto the soil were not consider. As a novelty, the
119 M4EKR version used in this work has been improved with the implementation of a kinetic
120 model of calcite dissolution/precipitation, extrapolable to any other mineral if present.

121 Appendix A presents the basic mathematical formulation employed in module M4EKR.
122 A more detailed description of the formulation can be found in literature [26].

123 **2.1. Calcite mass balance**

124 The evolution of the mass of calcite is obtained solving the mass balance equation

$$125 \quad \frac{dM_{calcite}}{dt} = R_{calcite} \quad \text{Eq. (1)}$$

126 where $M_{calcite}$ is the total mass of calcite (mol/m^3) and $R_{calcite}$ is the calcite
127 production/consumption rate ($\text{mol m}^{-3} \text{s}^{-1}$).

128 **2.2. Kinetic of calcite precipitation/dissolution**

129 The calcite production/consumption rate, $R_{calcite}$ ($\text{mol m}^{-3}\cdot\text{s}^{-1}$), has been estimated by a
130 general equation provided by Lasaga [41]:

$$131 \quad R_{calcite} = \pm k_{calcite} \cdot \phi \cdot Sr \cdot \rho_w \cdot S_{calcite} \cdot (1 - \Omega_{calcite}^\theta)^\eta \quad \text{Eq. (2)}$$

132 where $k_{calcite}$ is the kinetic constant ($\text{mol s}^{-1}\text{m}^{-2}$) of dissolution (negative value) or
133 precipitation (positive value) processes, $S_{calcite}$ is the reactive surface area ($\text{m}^2 \text{kg}_{\text{H}_2\text{O}}^{-1}$), and
134 $\Omega_{calcite}$ is the saturation ratio. Finally, θ and η are empirical parameters that determine the
135 relation between reaction rate and saturation ratio. The parameters selected for calcite
136 were $\theta = 0.5$ and $\eta = 2$ [42]. A reference value for $S_{calcite}$ was extracted from literature
137 [42]. The parameters ϕ , Sr and ρ_w correspond to the soil porosity, saturation degree and

138 water density. The saturation ratio ($\Omega_{calcite}$) is estimated using the Eq. (3) where $K_{s, calcite}$
 139 is the solubility constant of calcite [43] and IAP is the ionic activity product, calculated
 140 as indicated in Eq. (4) where $a_{Ca^{2+}}$ and $a_{CO_3^{2-}}$ are the activities of Ca^{2+} y CO_3^{2-}
 141 respectively.

$$142 \quad \Omega_{calcite} = \frac{IAP_{calcite}}{K_{s,calcite}} \quad \text{Eq. (3)}$$

$$143 \quad IAP_{calcite} = a_{Ca^{2+}} \cdot a_{CO_3^{2-}} \quad \text{Eq. (4)}$$

144 The kinetic constant ($k_{calcite}$) has been defined using a model that takes into account the
 145 influence additional mechanisms in the value of $k_{calcite}$. In this case, a dependence on the
 146 activity HCO_3^- has been observed for carbonate minerals such as calcite [42]. The
 147 expression for $k_{calcite}$ reads as follows

$$148 \quad k_{calcite} = k_{25}^{calcite} \cdot \exp \left[\frac{-E_a^{calcite}}{R} \left(\frac{1}{T} - \frac{1}{298.15} \right) \right] + k_{25}^{HCO_3^-} \cdot \exp \left[\frac{-E_a^{HCO_3^-}}{R} \left(\frac{1}{T} - \frac{1}{298.15} \right) \right] \cdot (a_{HCO_3^-})^{n_{HCO_3^-}}$$

149 Eq. (5)

150 where E_a^i and a_i are the activation energy (kJ mol^{-1}) and the activity, and k_{25}^i and n_i are a
 151 model parameters for the species considered, in this particular case the mineral considered
 152 (calcite) and HCO_3^- . Table 1 shows the numerical values of parameters for the kinetic
 153 model selected [42].

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159 **Table 1.** Kinetic model parameters from calcite

Parameters	Units	Values
$k_{25}^{calcite}$	$\text{mol m}^{-2} \text{s}^{-1}$	$1.8 \cdot 10^{-7}$
$k_{25}^{HCO_3^-}$	$\text{mol m}^{-2} \text{s}^{-1}$	$1.9 \cdot 10^{-3}$
$E_a^{calcite}$	kJ mol^{-1}	66
$E_a^{HCO_3^-}$	kJ mol^{-1}	67
$n_{HCO_3^-}$	-	1.63
θ	-	0.5
η	-	2
$S_{calcite}$	$\text{m}^2 \text{kg}_{\text{H}_2\text{O}}^{-1}$	1.54

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161 3. NUMERICAL IMPLEMENTATION

162 The M4EKR model was implemented in COMSOL Multiphysics [44], a partial
 163 differential equation solver developed in a multiphysics environment. COMSOL uses the
 164 finite element method with Lagrange multipliers. One of its most remarkable attributes is
 165 its versatility, since the user has great freedom to define the system of equations required
 166 to solve the selected case [45-47].

167 M4EKR solves transport problems coupled with chemical speciation using a monolithic
 168 approach, without using any other external software. This strategy optimises the resources
 169 and reduces the computational cost. This aspect is of high importance and provides a
 170 novelty to this work because it is common to use operator-splitting procedures for solving
 171 reactive transport problems [48-50]. Furthermore, the computational capacity of the
 172 M4EKR module goes beyond this kind of codes since it also allows to solve
 173 simultaneously the partial differential equations that define the transport and balance of
 174 mass of water, mass of dissolved species and electric charge, (See Appendix A, Eqs. A.1,
 175 A.9 and A.10 respectively), the ordinary differential equations used to solve the mass

176 balance in electrolyte wells, (Eq. A.11) and the algebraic equations from the
177 stoichiometric approach used in the chemical speciation.

178 **4. SIMULATION OF THE EKR PROCESS**

179 **4.1. Modelled configuration**

180 **4.1.1. Polluted Soil**

181 The electrokinetic remediation of a polluted soil with an organochlorine herbicide (20 mg
182 of clopyralid $\text{kg}_{\text{dry soil}}^{-1}$) was analysed. The selected soil is a calcareous soil, (calcite mass
183 fraction 10%). Table 2 presents the textural and mineralogical properties of the modelled
184 soil. The soil parameters are show in Table 3. Bulk density and gravimetric water content
185 are representative of the natural state of this soil [30].

186 **Table 2.** Textural and mineralogical properties of the modelled soil.

Mineralogical Analysis	
Mineral	%
Quartz	9
Feldspar	12
Calcite	10
Kaolinite	23
Smectite	25
Illite	21
Textural Parameters	
Size fraction	%
Sand	26.1
Silt	68.6
Clay	5.3

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192 **Table 3.** Modelled soil parameters.

Parameters	Description	Values	Units
α_{VG}	Parameter of the Van Genuchten[51] retention curve	0.0147	kPa ⁻¹
n_{VG}	Parameter of the Van Genuchten[51] retention curve	1.2593	-
m_{VG}	Parameter of the Van Genuchten[51] retention curve	0.2059	-
ϕ	Porosity	0.4681	-
K_{sat}^h	Saturated hydraulic permeability	2.03×10^{-10}	m s ⁻¹
K_{sat}^{eo}	Saturated electroosmotic permeability	2.4×10^{-9}	m ² V ⁻¹ s ⁻¹
ρ_s	Soil particle density	2681.5	kg m ⁻³
ρ_{bulk}	Bulk density	1894.4	kg m ⁻³
w	Moisture	0.32	kg _{water} kg _{dry soil} ⁻¹
δ_i^L	Longitudinal dispersivity of species i	0.01	m
τ	Vapour tortuosity	1.00	-

193

194 **4.1.2. Porewater**

195 The geochemical model implemented consists of 7 components and 20 secondary species.

196 The chemical reactions and thermodynamic properties (equilibrium constants [43], hard-
 197 core diameters and diffusion coefficients) are presented in Table 4. The activity
 198 coefficients of the species were estimated by WATEQ Debye-Hückel model [52].

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205 **Table 4.** Thermodynamic properties of the modelled geochemical systems

	Species	Reactions	log K_i^{eq} (25°C)	Hard core diameter / Å	$D_i^0 / \text{m}^2\text{s}^{-1}$
<i>Components</i>	Cl ⁻	Cl ⁻	0	3.6	2.03×10^{-9}
	H ₂ O	H ₂ O	0	3.4	5.27×10^{-9}
	H ⁺	H ⁺	0	4.1	9.31×10^{-9}
	CO ₃ ⁻²	CO ₃ ⁻²	0	4.7	9.55×10^{-10}
	Ca ⁺²	Ca ⁺²	0	5.7	7.93×10^{-10}
	Na ⁺	Na ⁺	0	4.1	1.33×10^{-9}
	ClOpy ⁻	ClOpy ⁻	0	3.6	$8.20 \times 10^{-10*}$
<i>Secondary species</i>	OH ⁻	H ₂ O ↔ OH ⁻ + H ⁺	-14	3.6	5.27×10^{-9}
	HCO ₃ ⁻	CO ₃ ⁻² + H ⁺ ↔ HCO ₃ ⁻	10.33	3.6	1.18×10^{-9}
	H ₂ CO ₃	CO ₃ ⁻² + 2H ⁺ ↔ H ₂ CO ₃	16.68	3.4	1.92×10^{-9}
	CaHCO ₃ ⁺	Ca ⁺² + CO ₃ ⁻² + H ⁺ ↔ CaHCO ₃ ⁺	11.43	4.1	5.06×10^{-10}
	CaCO ₃	Ca ⁺² + CO ₃ ⁻² ↔ CaCO ₃	3.22	3.4	4.46×10^{-10}
	Ca(OH) ⁺	Ca ⁺² + H ₂ O ↔ Ca(OH) ⁺ + H ⁺	-12.78	4.1	$2.13 \times 10^{-10*}$
	NaHCO ₃	Na ⁺ + HCO ₃ ⁻ ↔ NaHCO ₃	10.08	3.4	6.73×10^{-10}
	NaCO ₃ ⁻	Na ⁺ + CO ₃ ⁻² ↔ NaCO ₃ ⁻	1.27	3.6	5.85×10^{-10}
	NaOH	Na ⁺ + H ₂ O ↔ NaOH + H ⁺	-14.75	3.4	$1.89 \times 10^{-10*}$
	CaCl ⁺	Ca ⁺² + Cl ⁻ ↔ CaCl ⁺	-0.29	4.1	$2.13 \times 10^{-10*}$
	NaCl	Na ⁺ + Cl ⁻ ↔ NaCl	-0.5	3.4	$1.89 \times 10^{-10*}$
	CaCl ₂	Ca ⁺² + 2 Cl ⁻ ↔ CaCl ₂	-0.64	3.4	7.54×10^{-10}
	ClOpy ⁻	ClOpy ⁻ + H ⁺ ↔ ClOpy	2.32	3.4	$8.20 \times 10^{-10*}$
<i>Solid species</i>	Calcite	Calcite ↔ Ca ⁺² + CO ₃ ⁻²	-8.48**	-	-

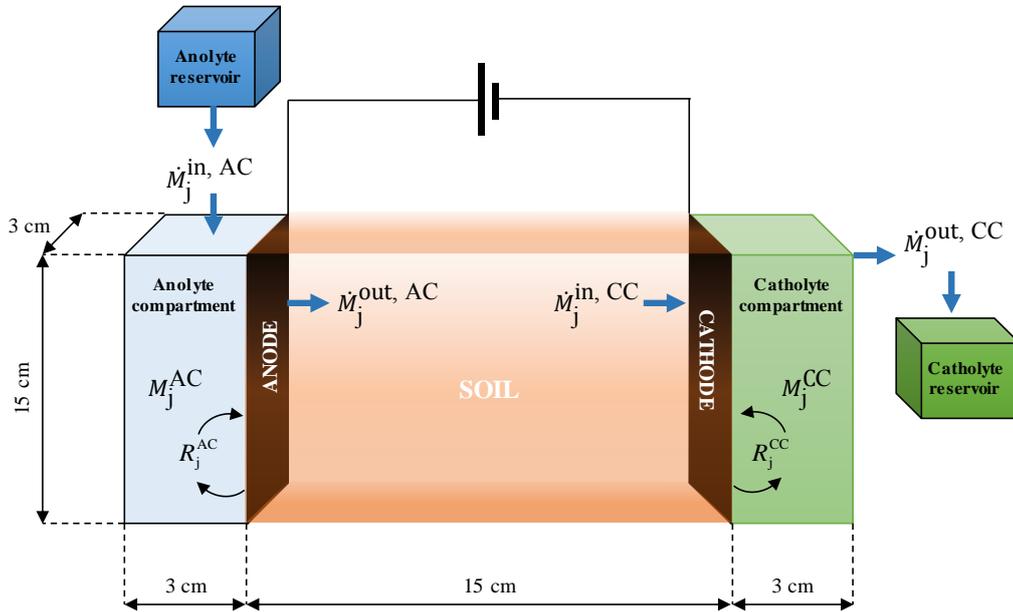
206 * Obtained by using Pikal's model [53].

207 ** Solubility product of calcite [43].

208 3.1.3. Modelled Experimental setup

209 The experimental setup modelled was a lab-scale EKR reactor. This setup is composed
 210 of an electrokinetic cell divided in three compartments: the anolyte compartment(AC),
 211 the catholyte compartment (CC), where electrodes are located and with dimensions
 212 15×3×3 cm, and a central compartment destined to place the polluted soil with a capacity
 213 of 675 ml (LWH: 15×15×3 cm). The level of electrolyte in compartments is constant
 214 during the tests by adding anolyte through gravity and extracting the catholyte by
 215 overflow of the CC. This modelled experimental setup has been outlined in detail in

216 preceding studies [26, 36, 37]. Fig. 1 shows a scheme of the conceptual model of the setup
 217 simulated.



218

219 **Figure 1.** Conceptual model of the one-dimensional domain proposed for simulation. Symbols
 220 described in Appendix A.

221 This setup defines a one-dimensional (1D) configuration that allows reducing the
 222 computational time and provides easily interpretable results needed to understand the
 223 EKR main trends.

224 3.1.4. Initial conditions

225 The initial mass fraction of calcite in the soil was 10%. The initial electrolytes (anolyte
 226 and catholyte) are an aqueous solution of CaCO_3 and NaCl equilibrated with the calcite
 227 present in the soil. The same solution with an amount of clopyralid equivalent to 20 mg
 228 $\text{kg}_{\text{dry soil}}^{-1}$ is considered as initial porewater. The electrolyte added to the AC during the
 229 EKR test is an aqueous NaCl solution with the same pH as the initial porewater. Table 5
 230 shows the chemical initial conditions of the different solutions used in the study.

231

232 **Table 5.** Chemical initial conditions.

	Total components concentration / m						pH	Ionic Strength / m
	H ⁺	Na ⁺	Ca ²⁺	CO ₃ ²⁻	Cl ⁻	Cl _{oxy} ^{anion}		
<i>Soil porewater</i>	3.41×10 ⁻³	3.58×10 ⁻³	2.49×10 ⁻³	3.12×10 ⁻³	5.36×10 ⁻³	3.67×10 ⁻⁴	7.25	1.10×10 ⁻²
<i>Initial electrolyte</i>	3.41×10 ⁻³	3.58×10 ⁻³	2.49×10 ⁻³	3.12×10 ⁻³	5.75×10 ⁻³	0	7.25	1.10×10 ⁻²
<i>Addition to anolyte reservoir</i>	-1.32×10 ⁻⁷	3.58×10 ⁻³	0	0	3.57×10 ⁻³	0	7.25	3.57×10 ⁻³

233

234 Initially, the electrolyte compartments are filled completely and they are open to
 235 atmosphere, therefore the initial liquid pressure, P_L , is equal to 100 kPa (atmosphere
 236 pressure). The initial condition was an electric potential gradient equal to zero.

237 3.1.5. Boundary conditions

238 To solve the problem presented in this work it is necessary to define three types of
 239 boundary conditions for each of the balances presented: electrical, hydraulic and mass of
 240 dissolved chemical species. For the first case, an average electrical potential gradient of
 241 1 V cm⁻¹ was applied. In the case of water mass balance, the boundary condition with the
 242 initial condition ($P_L = 100$ kPa) since the compartments are open to the atmosphere and
 243 maintain a constant level during all the cases evaluated. The boundary condition used in
 244 the mass balances of chemical species was imposed by the concentration of components
 245 in the electrolyte compartments. It is important to note that the definition of this type of
 246 boundary condition is obtained in a coupled form by means of the resolution of J-2
 247 ordinary differential equations defining the mass balance of the chemical components in
 248 the electrolyte wells. (See Appendix A, Eq. A.11).

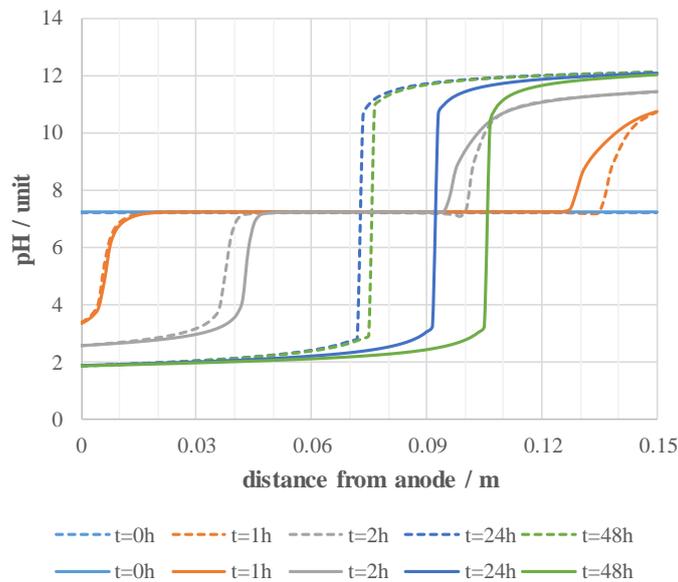
249

250 **3.2. Simulation of EKR tests**

251 **3.2.1. Effect of calcite dissolution-precipitation process in EKR behaviour.**

252 In this section, the importance of calcite dissolution-precipitation processes in the
253 performance of an EKR process has been evaluated. To this end, two EKR processes have
254 been studied: (i) the treatment of the described soil contaminated with clopyralid, in
255 which the dissolution-precipitation processes of this mineral have been taken into account
256 (Mineral Reaction test, MR-test) and (ii) an EKR process equivalent to the first case, in
257 which it has been assumed that there are no precipitation-dissolution processes of calcite
258 (no Mineral Reaction test, nMR-test). This strategy has been selected to determine the
259 effect of these reactions on the overall performance and behaviour of the selected EKR
260 process.

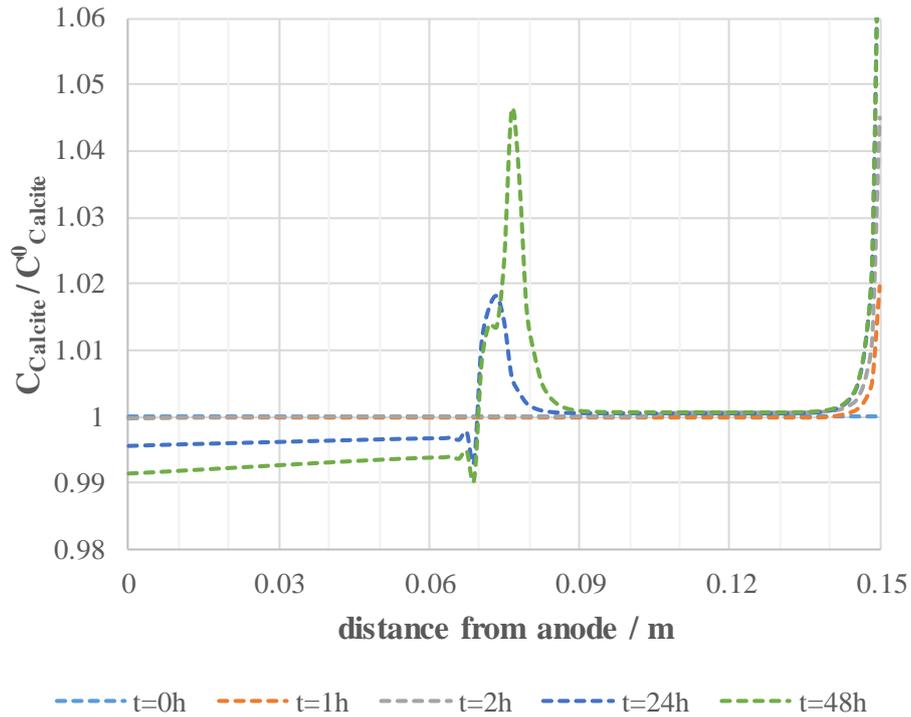
261 Figure 2 shows the spatial distribution of the pH in both tests for different observation
262 times (0, 1, 2, 24 and 48 h).



263

264 **Figure 2.** pH spatial distribution at selected times. MR-test, dashed line; nMR-test, solid line.

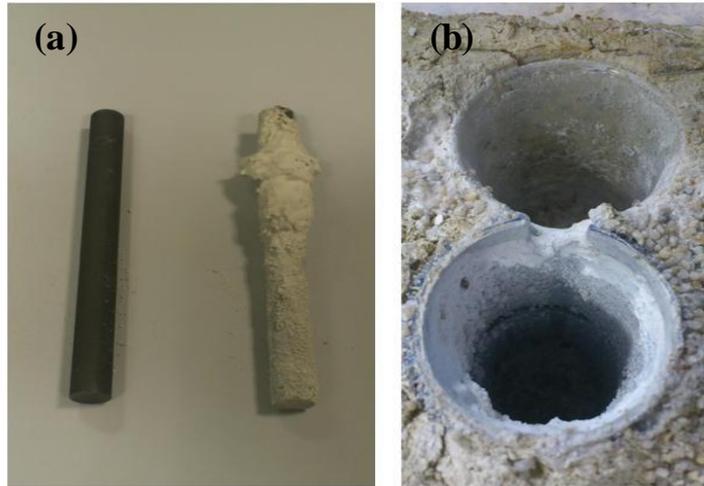
265 In both tests, the pH distribution observed in the domain is characteristic of EKR treated
266 soil. At short times (1 h), the generation of a pH gradient between the anodic (pH=3.5)
267 and cathodic (pH=10.5) zones due to the generation of H^+ in the AC and OH^- in the CC
268 through the electrolysis of water is already observed. The transport of these ionic species
269 through the soil generates an acidic pH front (to the cathode) and an alkaline pH front (to
270 the anode) that collide at a point where the pH gradient is maximum. The location of this
271 point will be very important in the discussion of results, the Maximum pH Gradient Point
272 (MpHGP). It can be seen how the velocity with which the pH fronts move is different in
273 the two numerical tests evaluated. The rate of advance of the acidic front is slower in the
274 case of MR-test. This decrease is more significant at higher test times. It can be seen that
275 after 48 hours of testing the acidic front has spread over 10.5 cm of the domain in the
276 nMR-test, and only 7.5 cm in the MR-test. This behaviour suggests that the MR-test
277 involves additional processes that act as a pH buffer such as the dissolution of calcite
278 [54]. The spatial distribution of the calcite content in the MR-test is shown in Fig. 3. It is
279 important to note that initially the calcite content is in equilibrium with the calcium
280 carbonate concentration.



281

282 **Figure 3.** Spatial distribution of normalized calcite content at selected times in MR-test.

283 On the other hand, when the content of calcite in the soil is analysed, three areas with
 284 different behaviour can be identified. In the zone with acidic pH, between AC and
 285 MpHGP, a decrease in the initial calcite content of the soil is observed, accentuating in
 286 the zone immediately adjacent to MpHGP. The calcite dissolution process is active under
 287 these conditions, and this trend is most noticeable at higher test times (24 and 48 h)
 288 because these processes have been simulated by means of a chemical kinetic process and
 289 not by a chemical equilibrium. Meanwhile, in the soil-CC interface zone (where the
 290 cathode is located) a sudden increase in the calcite content generated by the precipitation
 291 can be seen. This behaviour has been observed (Fig. 4) in previous experimental work
 292 [55].

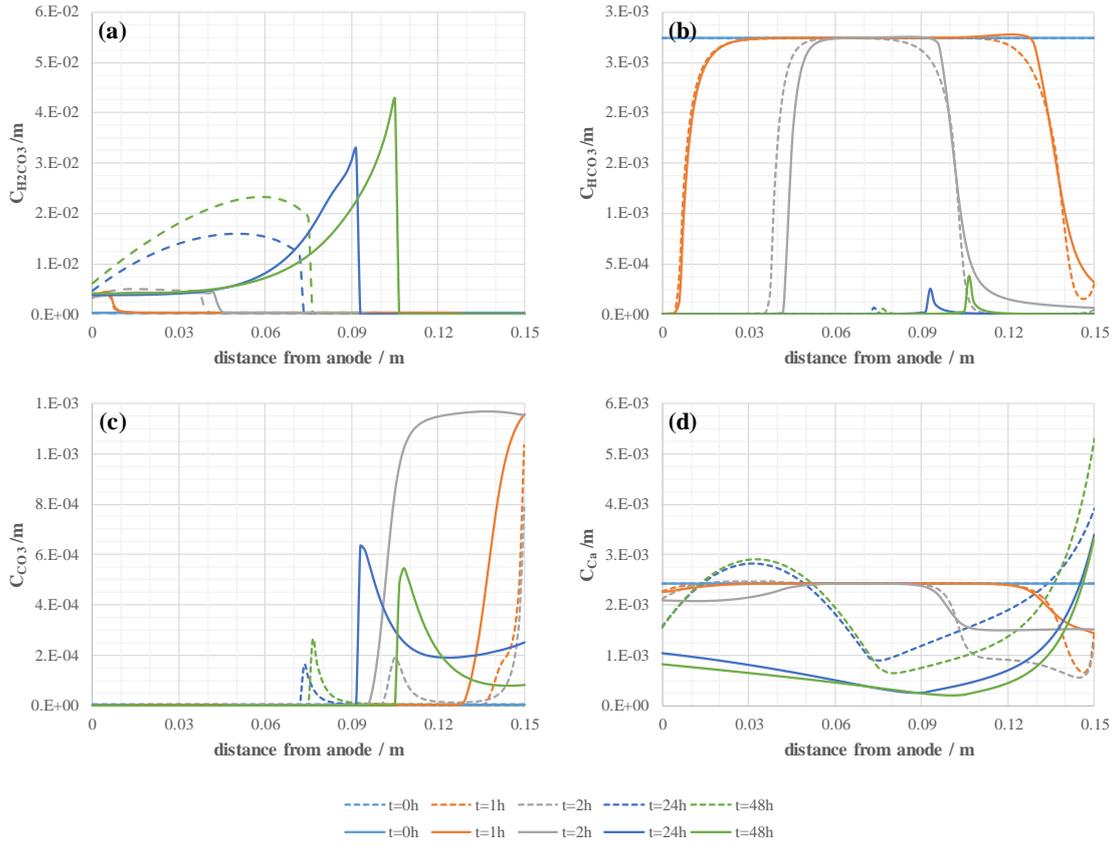


293

294 **Figure 4.** Evidences of mineral precipitation on (a) cathode before EKR and (b) CC.

295 The third zone which was identified corresponds to the domain close to the MpHGP
296 where the acidic and alkaline pH fronts (maximum pH gradient) are located. In this area
297 there is an accumulation of calcite. In both the second and third zones, the observed
298 calcite precipitation is limited only to very localize areas and not to the entire soil domain
299 with basic pH, as would be expected. In order to understand more precisely the processes
300 of calcite dissolution-precipitation observed in Figure 3, it is necessary to analyze the
301 chemical speciation of the carbonate system. For this purpose, the spatial distribution of
302 the concentration of the carbonate species ($\text{H}_2\text{CO}_3^\circ$, HCO_3^- , CO_3^{2-} y Ca^{2+}) is shown in Fig.
303 5.

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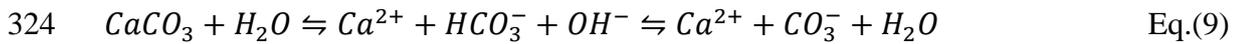


305

306 **Figure 5.** Spatial distribution of species concentration at selected times. MR-test, dashed line;
 307 nMR-test, solid line. (a) $H_2CO_3^\circ$, (b) HCO_3^- , (c) CO_3^{2-} and (d) Ca^{2+} .

308 Fig. 5a shows the distribution of $H_2CO_3^\circ$ (species including the sum of the concentrations
 309 of H_2CO_3 and dissolved CO_2) in the domain. A common trend is observed in both
 310 numerical tests, the accumulation of this species in the post-MpHGP area. This can be
 311 explained taking into account that this species is only present in systems with pH highly
 312 acid. In addition, taking into account the transport features, this species has a neutral
 313 charge, so the predominant transport mechanisms are the advective flux (hydraulic and
 314 electro-osmotic flow) and the diffusive/dispersive processes with a net flux from anode
 315 to cathode [26]. However, if we compare the spatial distributions at the same observation
 316 times, we can see how there are notable differences between the two tests analysed. In
 317 the MR-test it is observed the $H_2CO_3^\circ$ is present in a smaller volume of soil, which
 318 indicates that there is a relationship between the concentration of $H_2CO_3^\circ$ and the process

319 of dissolution of calcite. As Plummer demonstrated [39], this process takes place through
320 simultaneous reactions (See Eq.5-9).



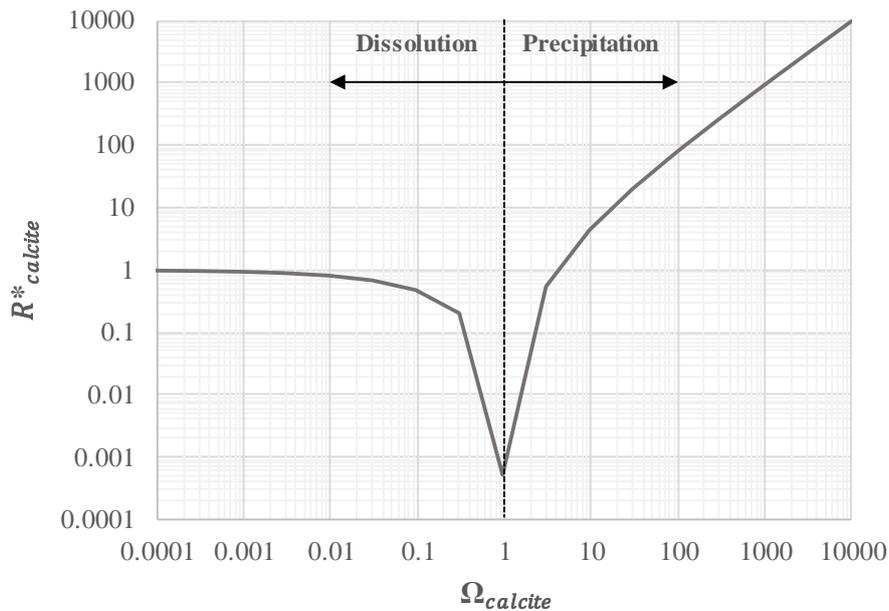
325 The progression of these reactions depends directly on the pH of the system and therefore
326 on the chemical speciation of the carbonate system. [56]. In systems with strongly acidic
327 pHs, the Eq.6 and Eq.7 reactions become important. Taking into account this reactivity,
328 the consumption of H^+ produced by the water electrolysis could justify the slowing down
329 observed in the MR-test of the acidic front (See Fig. 2), the reduction of the calcite content
330 by dissolution (See Fig. 3) and the increase in the concentration of H_2CO_3 and
331 accumulation in the area after the acidic front (See Fig. 5a). In addition, the decrease of
332 the calcite content in the left side of the MpHGP point could be explained if it is known
333 that the predominant reaction of the calcite solution is the Eq.6 reaction and if it is taken
334 into account that for pH greater than 5 and with a partial pressure of CO_2 greater than 0.1
335 atm, Eq.8 becomes important. With the contribution of the latter reaction to the general
336 process of dissolution, the change in shape of the spatial distribution of H_2CO_3 mentioned
337 above can be satisfactorily described.

338 The second carbonate species analysed was the bicarbonate ion, HCO_3^- (Fig. 5b). This
339 species exists at pH values between 6.3 and 10.25. For this reason, HCO_3^- is present in
340 significant amounts only in the initial stage of the test (1 and 2 h), when the pH fronts are
341 not very abrupt. The ion carbonate, CO_3^{2-} , exhibits the opposite behavior to $H_2CO_3^{\circ}$. The

342 CO_3^{-2} is predominantly presents in the right side of the MpHGP and has negative charge
 343 as well as HCO_3^- . Both species are predominantly transported by electromigration to the
 344 anode. These clarifications could explain the spatial distribution of the concentration of
 345 these species and the accumulation observed behind the MpHGP (See Fig. 5b and 5c).
 346 Quantitatively, the MR-test presents a lower accumulation of both CO_3^{-2} and HCO_3^- in the
 347 area near the MpHGP since these species are, together with Ca^{2+} (See reaction R.9), the
 348 drivers of the generation of calcite by precipitation (See Fig.3). This behaviour is also
 349 observed in the area of the soil-CC interface, even at short test times (1 and 2h). This
 350 suggests that the precipitation process is faster than the dissolution process, as there is no
 351 decrease in calcite content at those times. In order to check this fact, the dimensionless
 352 reaction rate $R_{calcite}^*$ has been evaluated as a function of the saturation ratio, $\Omega_{calcite}$ using
 353 the expression (10) and giving the results in Figure 6.

354
$$R_{calcite}^* = (1 - \Omega_{calcite}^\theta)^\eta \quad \text{Eq. (10)}$$

355 and giving the results in Figure 6.



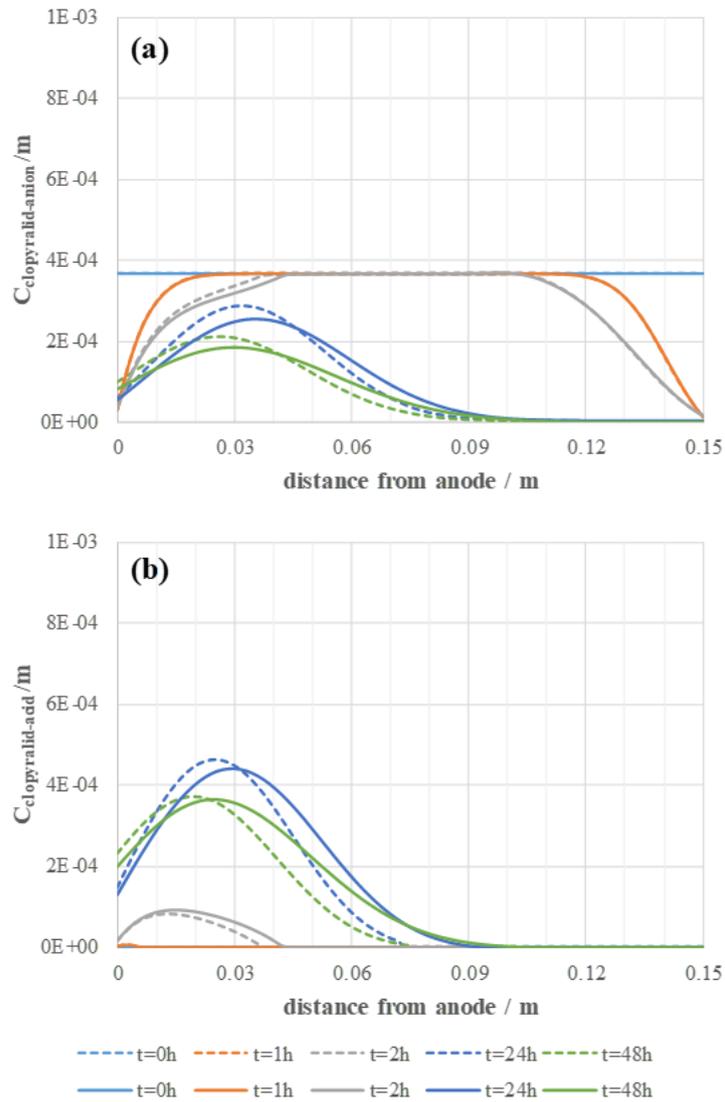
356

357 **Figure 6.** Influence of saturation ratio in dimensionless reaction rate.

358 Starting from a system in equilibrium ($\Omega_{calcite}=1$), it can be observed that a decrease of
359 $\Omega_{calcite}$ activates the process of dissolution of calcite. The rate of the calcite dissolution
360 process has an increasing trend corresponding to the decrease of the $\Omega_{calcite}$. This
361 evolution continues until it reaches a point where the reaction rate is stabilized at a
362 constant value. This implies that from this point on, the dependence of the reaction rate
363 of the dissolution process with $\Omega_{calcite}$ is negligible. Meanwhile, the precipitation
364 process occurs when $\Omega_{calcite}$ is greater than 1. The evolution of the precipitation reaction
365 rate as a function of $\Omega_{calcite}$ is entirely different. A roughly double log-linear increase is
366 observed with the increase in the saturation ratio. This indicates that the precipitation rate
367 is proportional to the concentration of species involved in that reaction (See reaction
368 Eq.8). In this line, from a quantitative point of view, it can be observed how changes in
369 an order of magnitude in $\Omega_{calcite}$ are translated into much more significant changes in
370 the reaction rate in the precipitation process. This would explain the rapid calcite
371 deposition shown in Fig. 3. Finally, the spatial distribution of Ca^{2+} ion is analysed (Fig.
372 5d). Both tests show an accumulation of this species in the area of the CC, mainly due to
373 the contribution of the electromigration process to the net transport of this ion [26]. The
374 most noticeable difference between the two tests analysed is the increase in concentration
375 observed in the acidic zone (after MpHGP) in the MR-test. This trend is according with
376 the process of calcite dissolution observed in this zone.

377 Subsequently, the influence of the buffer effect generated by the dissolution of calcite on
378 the speciation and transport of clopyralid was analysed. Two species have been
379 considered, acid and anionic.

380



381

382 **Figure 7.** Spatial distribution of the total concentration of (a) Clopyralid_{aniion} species and (b)
 383 Clopyralid_{acid} species at selected times. MR-test, dashed line; nMR-test, solid line.

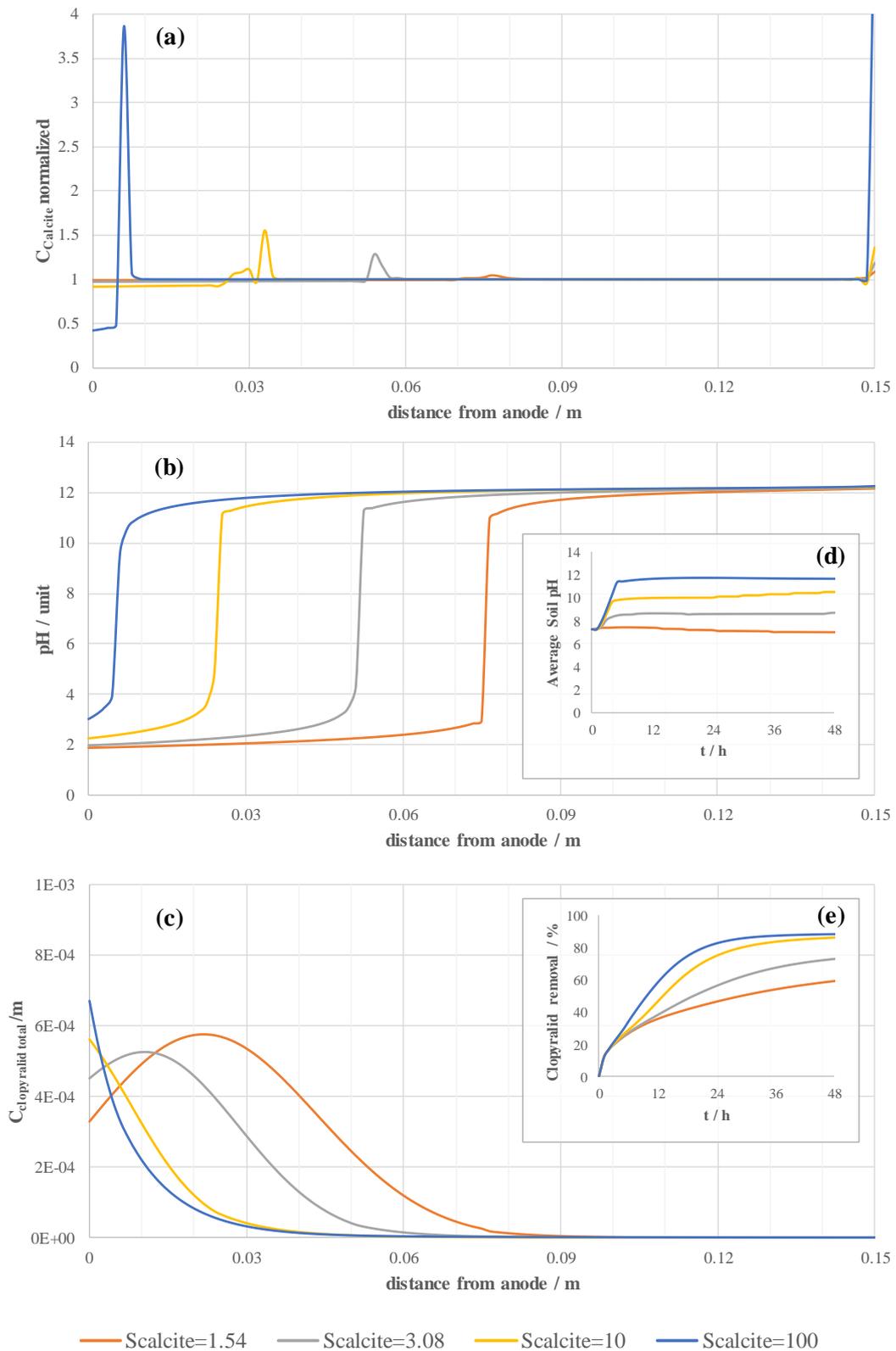
384

385 The pH of the soil pore water affects the chemical speciation of the substances contained
 386 in the soil [22, 57], so acidic species are present in the low pH soil zone, generally between
 387 CA and MpHGP in EKR processes. In the case studied, Clopyralid_{acid} species are
 388 concentrated in this zone, favouring the transport of these species to the anodic
 389 compartment mainly by diffusion [36]. Furthermore, Clopyralid_{aniion} species, in
 390 equilibrium with acidic species, also accumulate in this zone and is susceptible to being

391 moved towards the anodic compartment mainly by electromigration [36]. For these
392 reasons, the performance of the EKR treatment of soils with this type of pollutant (acid
393 pollutant with anionic dissociation) improves when the volume of the acid soil zone is
394 reduced as has been proven in previous works [36]. The buffer effect generated by the
395 dissolution of calcite produces a reduction of the acidic zone. This is how it can be seen
396 in Fig 7. Fig 7 shows in those locations where an accumulation of contaminant is observed
397 (in one of the two species considered) it is always shown to a lesser extent in the MR-
398 test.

399 **3.2.2. Sensitivity analysis of reactive surface area in EKR behaviour.**

400 The buffer capacity of the soil depends directly on the precipitation/dissolution kinetics
401 of calcite, however, as previously discussed in the introduction, it is considerably
402 complex to accurately estimate the kinetics due to the high uncertainty associated with
403 the reactive surface area [39]. For this reason, a sensitivity analysis of the simulated EKR
404 process has been carried out for different values of this parameter. For this purpose, three
405 additional MR-tests have been simulated in which a reactive surface area ($S_{calcite}$) value
406 of 3.08, 10 y 100 $\text{m}^2 \text{kg}_{\text{H}_2\text{O}}^{-1}$ (compared to the 1.54 used in the case analysed in section
407 3.2.1), has been assumed. The main results obtained in this study are shown in Fig. 8.



408

409 **Figure 8.** Sensitivity analysis of reactive surface area. Spatial distribution of the (a) of normalized
 410 calcite content, (b) pH and (c) total concentration of Clopyralid species in soil at time test of 48
 411 h. Dynamic response of the (d) average soil pH and (e) Clopyralid removal.

412 From a kinetic point of view, an increase in the value of $S_{calcite}$ translates into an increase
413 in the rate at which the dissolution and precipitation processes of calcite take place.
414 However, qualitatively, no significant influence is found. This can be verified by
415 analysing Fig. 8a. The behaviour observed in the cases studied is common; the same three
416 previously identified areas with active calcite reactivity are shown (Fig.3): (i) generalised
417 dissolution in the acidic zone and punctual precipitation close to (ii) MpHGP and (iii)
418 cathode. The extent to which these reactions develop is directly related to the $S_{calcite}$ value
419 used.

420 Additionally, if the calcite dissolution process is specifically analyzed for a value of $S_{calcite}$
421 =100, a decrease in the mineral content of the soil is observed and therefore the
422 consumption of H^+ is more pronounced This statement is in accordance with the pH
423 profiles obtained in the MR-tests (Fig 8.b). It can be seen that as $S_{calcite}$, increases, and
424 therefore the kinetics of calcite dissolution, the speed of the acidic front slows down. In
425 order to be able to evaluate the buffer capacity of the soil in global terms, the temporal
426 evolution of the average pH has been analysed (Fig. 8d). The results obtained confirm
427 that an increase in $S_{calcite}$ causes a significant improvement in soil buffer capacity. In the
428 MR-test with $S_{calcite}=100$ an average pH value of 11.6 is obtained in 5 h of test, which
429 means that in this case the acidic front is totally neutralized.

430 As discussed in the previous section, the pH of the porewater determines chemical
431 speciation and, consequently, the distribution of chemical species within the soil.
432 Therefore, the spatial distribution of the pollutant has been analysed (Fig. 8c). The
433 neutralization of the acidic front favours the accumulation of the contaminant in the area
434 close to the CA and therefore the transport to this compartment. An increase in the soil
435 buffer capacity produces an improvement in the performance of the EKR process,
436 reaching an elimination rate of close to 90% (Fig. 8e). Although these results cannot be

437 taken valid in any configuration and for any soil, the information obtained in this
438 sensitivity study is valid to show how the soil buffer capacity strongly influences the net
439 performance of an EKR treatment.

440 **4. CONCLUSIONS**

441 This paper presents a numerical study of the processes of dissolution-precipitation of
442 calcite in soil contaminated with clopyralid (an herbicide of the organochlorine acid type)
443 by using an EKR treatment. The influence of the soil buffer capacity (associated with the
444 dissolution of calcite) on the overall performance of the remediation process is also
445 studied. In addition, experimentally observed behaviours, such as rapid calcite
446 precipitation in the cathode electrode and the compartment where it is located, have been
447 identified in the simulations. Thanks to the speciation of the carbonate system, it has been
448 possible to improve the relative understanding of the reactive transport of these chemical
449 species.

450 In addition, the results obtained have highlighted the importance of considering the
451 processes of dissolution-precipitation in an EKR treatment, since the distribution of the
452 soil pH depends to a large extent on them. The pH of the soil porewater is a fundamental
453 variable in EKR treatments as it directly influences the speciation of the contaminants
454 (and other chemical species present) and indirectly determines the transport mechanisms
455 involved in the elimination of these contaminants. In the cases evaluated, it has been
456 proven that the buffer effect produces improvements in the overall performance of the
457 EKR process. The reduction of the advance of the acidic front, it favours the accumulation
458 of polluting species in the vicinity of the anodic compartment and can be eliminated by
459 means of anionic species by electromigration and of the acidic species by diffusive flux.

460 All these conclusions have been confirmed after a sensitivity analysis in which it has been
461 evaluated how changes in the kinetics of the calcite dissolution-precipitation process
462 abruptly modify the buffer capacity of the soil and therefore strongly influence the
463 performance of the treatment. Therefore, it is very important to define accurately the
464 kinetic parameters (specifically, the reactive surface area) since the interpretation of the
465 results as well as the previous design of experimental setups depends on their numerical
466 values.

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477

478

479

480 **APPENDIX A**

481 **A.1. Water mass balance**

482 Eq. (A.1) show the expression that defines the water mass balance:

483
$$\frac{\partial m_w}{\partial t} + \nabla \cdot \mathbf{l}_w = 0 \quad \text{Eq. (A.1)}$$

484 where m_w is the mass of water per unit total volume (kg m^{-3}), and it is defined as Eq (A.2),

485 $\nabla \cdot$ is the divergence operator, and \mathbf{l}_w is the mass flux of water ($\text{kg m}^{-2} \text{s}^{-1}$).

486
$$m_w = \phi \cdot Sr \cdot \rho_w + \phi \cdot (1 - Sr) \cdot \rho_v \quad \text{Eq. (A.2)}$$

487 where ϕ is the soil porosity, Sr is the degree of saturation of the soil defined by van

488 Genuchten model [58], ρ_w and ρ_v are the liquid and vapour water density (kg m^{-3})

489 respectively. Vapour water density is estimated using the psychrometric equation, Eq

490 (A.3) [59]:

491
$$\rho_v = \rho_v^0 \cdot \exp\left(\frac{M_w \cdot s}{\rho_w \cdot R \cdot T}\right) \quad \text{Eq. (A.3)}$$

492 where M_w is the water molar mass, T is absolute temperature and, s is the matric suction,

493 assumed equal to the capillary one (osmotic suction is not considered).

494 Eq. (A.4) shows the expression to calculate the flux term \mathbf{l}_w :

495
$$\mathbf{l}_w = \mathbf{l}_w^h + \mathbf{l}_w^{\text{eo}} = \rho_w \cdot (\mathbf{q}_w^h + \mathbf{q}_w^{\text{eo}}) = \rho_w \cdot \mathbf{q}_w \quad \text{Eq. (A.4)}$$

496 Where \mathbf{l}_w^h is the Darcian flux, and \mathbf{l}_w^{eo} is the electroosmotic flux, estimated with the semi-

497 empirical model of Helmholtz-Smoluchowski [60]. The hydraulic and electroosmotic

498 volumetric fluxes are presented as \mathbf{q}_w^h and \mathbf{q}_w^{eo} (See Eq. (A.5) and (A.7) respectively).

499
$$\mathbf{q}_w^h = -K_e^h \cdot (\nabla P_L + g \cdot \rho_w \cdot \nabla z) \quad \text{Eq. (A.5)}$$

500 $K_e^h = K_{sat}^h \cdot k_{rel} = K_{sat}^h \cdot Sr^3$ Eq. (A.6)

501 $\mathbf{q}_w^{eo} = -K_e^{eo} \cdot \nabla E$ Eq. (A.7)

502 $K_e^{eo} = K_{sat}^{eo} \cdot k_{rel} = K_{sat}^{eo} \cdot Sr^3$ Eq. (A.8)

503 where K_e^h is the effective hydraulic permeability, which is calculated by Eq (A6), K_{sat}^h
 504 is the saturated permeability and k_{rel} is relative permeability. K_e^{eo} is the effective electro-
 505 osmotic permeability, which is estimated by Eq (A8), and K_{sat}^{eo} is the saturated electro-
 506 osmotic permeability. Both relative permeability are calculated using a model Brooks and
 507 Corey [61] with an exponent of 3. Additionally, ∇ is the gradient differential operator, g
 508 is gravity, z is the vertical coordinate, E is the electric potential.

509 **A.2. Reactive transport**

510 The geochemical model implemented in M4EKR module contemplates J components
 511 which is capable of generating a sum of N chemical species for the chemical reactions
 512 system considered (See Table 4) [62]. Thus, to determine the total chemical composition
 513 of the system, it is necessary to calculate the total mass of each component j ($j=1\dots J$)
 514 using a general mass balance equation, Eq (A.9). However, it is not required to solve J
 515 mass balances, only $J-2$ partial differential equations are necessary to compute, due to the
 516 mass of one of the components is obtained applying the electroneutrality condition and
 517 water mass is calculated by Eq. (A.1) .

518 $\frac{\partial m_j}{\partial t} + \nabla \cdot \mathbf{I}_j = R_j$ Eq. (A.9)

519 where m_j is the total mass of component j per unit total volume (mol m^{-3}), \mathbf{I}_j is its total
 520 molar flux ($\text{mol m}^{-2} \text{s}^{-1}$) which is defined as the sum of the contributions of four transport
 521 processes: advective transport generated by (i) hydraulic, (ii) electroosmotic and (iii)

522 electromigration fluxes, and (iv) Fickian diffusive-dispersive transport; and R_j is its rate
 523 of production or consumption ($\text{mol m}^{-3} \text{s}^{-1}$). Only $J-2$ partial differential equations are
 524 required to be resolved, due to the total mass of one of the components is computed after
 525 electroneutrality and the total mass of water is calculated by Eq. (A.1).

526 The resolution of the chemical speciation has been obtained applying a classical
 527 stoichiometric approach based on to solve a system of mass-balance and mass-action
 528 equations [63]. The general algorithm of this method has been described in detail in
 529 literature [52, 62, 64, 65], even so the explicit formulation utilized in M4EKR can be seen
 530 in previous works [26, 37].

531 **A.3. Electric charge balance**

532 The balance equation of the total electric charge is defined by the expression Eq. (A.10)
 533 assuming that the electroneutrality condition is fulfilled throughout the entire domain and
 534 the system the system is not able to accumulate electric charge.

$$535 \quad \nabla \cdot \mathbf{i} = 0 \quad \text{Eq. (A.10)}$$

536 where \mathbf{i} is the total current density (A m^{-2}), which is estimated according Ohm's law [26,
 537 66]. Rhoades [67] and Appelo [68] approaches are applying to calculated the apparent
 538 electrical conductivity of the soil and electrical conductivity of pore water respectively.
 539 The equations of these methods can be found in López-Vizcaíno et al. ([26]).

540 **A.4. Mass balance in electrolyte wells**

541 The temporal evolution of the mass of the J components is obtained solving $J-2$ balance
 542 equations, Eq. (A.11) in each electrolyte compartment. The “ideal continuous stirring
 543 tank reactors” assumption is supposed in both electrolyte compartments.

$$544 \quad \frac{dM_j^*}{dt} = \dot{M}_j^{in,*} - \dot{M}_j^{out,*} + R_j^* \quad \text{Eq. (A.11)}$$

545 where M_j^* is the total mass of component j (mol) and \dot{M}_j denotes the mass flow of
546 component j (mol s⁻¹), being an input flow (superscript *in*) or output flow (superscript
547 *out*). Additionally, R_j^* is the production/consumption/ term (mol s⁻¹) linked to the
548 electrochemical reactions (only water electrolysis). The superscript * indicates the kind
549 of electrolyte compartment, which is AC or CC (anolyte or catholyte compartment
550 respectively). The equations employed to estimate these variables can be found in a
551 previous work [26, 36, 37].

552

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